Detailed chemical characterization of unresolved complex mixtures in atmospheric organics: Insights into emission sources, atmospheric processing, and secondary organic aerosol formation

Arthur W. H. Chan, ^{1,2} Gabriel Isaacman, ¹ Kevin R. Wilson, ³ David R. Worton, ^{1,4} Christopher R. Ruehl, ^{1,3} Theodora Nah, ⁵ Drew R. Gentner, ⁶ Timothy R. Dallmann, ⁶ Thomas W. Kirchstetter, ^{6,7} Robert A. Harley, ^{6,7} Jessica B. Gilman, ^{8,9} William C. Kuster, ⁹ Joost A. de Gouw, ^{8,9} John H. Offenberg, ¹⁰ Tadeusz E. Kleindienst, ¹⁰ Ying H. Lin, ¹¹ Caitlin L. Rubitschun, ¹¹ Jason D. Surratt, ¹¹ Patrick L. Hayes, ^{8,12} Jose L. Jimenez, ^{8,12} and Allen H. Goldstein ^{1,6,7}

Received 13 February 2013; revised 28 May 2013; accepted 29 May 2013.

[1] Recent studies suggest that semivolatile organic compounds (SVOCs) are important precursors to secondary organic aerosol (SOA) in urban atmospheres. However, knowledge of the chemical composition of SVOCs is limited by current analytical techniques, which are typically unable to resolve a large number of constitutional isomers. Using a combination of gas chromatography and soft photoionization mass spectrometry, we characterize the unresolved complex mixture (UCM) of semivolatile aliphatic hydrocarbons observed in Pasadena, California (~16 km NE of downtown Los Angeles), and Bakersfield, California, during the California Research at the Nexus of Air Quality and Climate Change 2010. To the authors' knowledge, this work represents the most detailed characterization of the UCM in atmospheric samples to date. Knowledge of molecular structures, including carbon number, alkyl branching, and number of rings, provides important constraints on the rate of atmospheric processing, as the relative amounts of branched and linear alkanes are shown to be a function of integrated exposure to hydroxyl radicals. Emissions of semivolatile branched alkanes from fossil fuel-related sources are up to an order of magnitude higher than those of linear alkanes, and the gas-phase OH rate constants of branched alkanes are ~30% higher than their linear isomers. Based on a box model considering gas/particle partitioning, emissions, and reaction rates, semivolatile branched alkanes are expected to play a more important role than linear alkanes in the photooxidation of the UCM and subsequent transformations into SOA. Detailed speciation of semivolatile compounds therefore provides essential understanding of SOA sources and formation processes in urban areas.

Citation: Chan, A. W. H., et al. (2013), Detailed chemical characterization of unresolved complex mixtures in atmospheric organics: Insights into emission sources, atmospheric processing, and secondary organic aerosol formation, *J. Geophys. Res. Atmos.*, 118, doi:10.1002/jgrd.50533.

Additional supporting information may be found in the online version of this article.

¹Department of Environmental Science, Policy, and Management, University of California, Berkeley, California, USA.

²Department of Chemical Engineering and Applied Chemistry, University of Toronto, Toronto, Ontario, Canada.

³Chemical Sciences Division, Lawrence Berkeley National Laboratory, Berkeley, California, USA.

⁴Aerosol Dynamics Inc., Berkeley, California, USA.

Corresponding author: A. W. H. Chan, Department of Chemical Engineering and Applied Chemistry, University of Toronto, 200 College St., Toronto, ON M5S 3E5, Canada. (arthurwh.chan@utoronto.ca)

^{©2013.} American Geophysical Union. All Rights Reserved. 2169-897X/13/10.1002/jgrd.50533

⁵Department of Chemistry, University of California, Berkeley, California, USA.

⁶Department of Civil and Environmental Engineering, University of California, Berkeley, California, USA.

⁷Environmental Energy Technologies Division, Lawrence Berkeley National Laboratory, Berkeley, California, USA.

⁸Cooperative Institute for Research in the Environmental Sciences, University of Colorado Boulder, Boulder, Colorado, USA.

⁹Chemical Sciences Division, NOAA, Boulder, Colorado, USA.

¹⁰National Exposure Laboratory, Office of Research and Development, U.S. Environmental Protection Agency, Research Triangle Park, North Carolina, USA.

¹¹Department of Environmental Sciences and Engineering, Gillings School of Global Public Health, University of North Carolina at Chapel Hill, Chapel Hill, North Carolina, USA.

¹²Department of Chemistry and Biochemistry, University of Colorado Boulder, Boulder, Colorado, USA.

1. Introduction

- [2] Organic matter represents a major fraction of ambient aerosol. A large fraction of organic aerosol (OA) is formed from oxidation of gas-phase hydrocarbons and is known as secondary organic aerosol (SOA) [Zhang et al., 2007]. However, current atmospheric models typically underestimate the budget of SOA, highlighting a lack of understanding in the sources and transformation processes of organic compounds in the atmosphere [Goldstein and Galbally, 2007; Hallquist et al., 2009]. Recently, it has been proposed that oxidation of semivolatile organic compounds (SVOCs) is a major source of SOA, particularly in urban atmospheres [Robinson et al., 2007; Pye and Seinfeld, 2010]. SVOCs are defined as compounds which have effective saturation concentrations (C^*) between 0.1 and 10 μ g m⁻³. Owing to the low vapor pressures of SVOCs and of their subsequent reaction products, oxidation of SVOCs is expected to lead to significantly higher SOA yields than more volatile precursors and, despite lower emissions, could dominate SOA formation in urban areas. Models that incorporate transformations of SVOCs to SOA yield better predictions of the total amount of SOA and its seasonal variability [Robinson et al., 2007; Hodzic et al., 2010]. However, the models fail to capture the chemical properties of the observed SOA and underestimate O/C ratios [Hodzic et al., 2010], indicating that the atmospheric chemistry of SVOCs leading to SOA formation is poorly understood. The relative oxidation rates of SVOCs and fragmentation/functionalization branching ratios, the two key factors that determine the amount and properties of SOA formed, are currently extrapolated from those of smaller carbon numbers.
- [3] Detailed knowledge of the identities and chemistry of SVOCs has been elusive, as the chemical composition is highly complex and current analytical techniques are unable to separate and identify these compounds. While gas chromatography-mass spectrometry (GC/MS) has been the most common method for speciating these compounds, atmospheric sources often emit a large number of compounds, which cannot be fully separated by traditional GC methods. As a result, many atmospheric samples contain a large unresolved peak in chromatographic analyses, often referred to as the unresolved complex mixture (UCM) [Schauer et al., 1999]. The UCM can account for more than 80% of semivolatile emissions from diesel [Schauer et al., 1999] and gasoline engines [Schauer et al., 2002], representing a major fraction of SVOCs in urban areas [Williams et al., 2010b]. Furthermore, the UCM is ubiquitous in environmental chemistry, often found in samples associated with fossil fuel use [Frysinger et al., 2003; Nelson et al., 2006; Ventura et al., 2008].
- [4] The UCM is thought to contain a large number of constitutional isomers, most of which are linear (also termed straight-chained or normal), branched, and/or cyclic alkanes [Mao et al., 2009]. In contrast, aromatic compounds comprise only a minor fraction of the UCM [Van Deursen et al., 2000]. Owing to the challenges in speciating alkane isomers, SVOCs have typically been classified by volatility only, inferred from thermodenuder measurements or gas chromatography retention times [Grieshop et al., 2009; Presto et al., 2012], which then serve as inputs into the aforementioned models [Hodzic et al., 2010; Pye and Seinfeld, 2010]. Information about

- molecular structures is generally overlooked. However, the number of rings and alkyl branches in alkanes strongly affects their oxidation chemistry and their SOA yields [Lim and Ziemann, 2009; Tkacik et al., 2012]. Molecular structure also plays an important role in other degradation pathways of environmentally relevant complex mixtures [Nelson et al., 2006]. As a result, knowledge of molecular structure is crucial to understanding the sources and environmental fate of these hydrocarbons.
- [5] The main objective of this work is to characterize the molecular composition of ambient UCM, classify the compounds by molecular structure for use in future models of urban atmospheres, and understand the atmospheric processing of SVOCs leading to SOA formation. Previously, we have demonstrated that constitutional isomers present in diesel fuel can be characterized using gas chromatography coupled to vacuum ultraviolet ionization mass spectrometry (GC/VUV-MS) [Isaacman et al., 2012b]. We combine this technique with comprehensive two-dimensional gas chromatography (GC × GC/VUV-HRTOFMS) to analyze the composition of ambient SVOCs observed during the California Research at the Nexus of Air Quality and Climate Change (CalNex) campaign in summer 2010. The samples were collected at two urban sites in California (Pasadena and Bakersfield). Using detailed speciated measurements of linear, branched, and cyclic alkanes in the UCM, we constrain their relative oxidation rates and provide valuable insights into their sources, processing, and relative contributions to SOA formation. The utility of the analytical technique in resolving aliphatic, aromatic, and oxygenated species observed in the ambient atmosphere is also discussed.

2. Methods

2.1. Sampling Locations

[6] The samples described in this paper were collected at two urban sites in California (Pasadena and Bakersfield) as part of the CalNex field campaign in the summer of 2010 [Ryerson et al., 2013]. Pasadena sampling was conducted on the campus of the California Institute of Technology, ~16 km NE of downtown Los Angeles (hereafter referred to as the LA site). Two PM_{2.5} samples from the LA site, collected on Sunday, 30 May (3:00–6:00 P.M. local time), and Saturday, 5 June (midnight to 11 P.M. local time), were selected for analysis using VUV ionization. Only two samples were chosen for VUV analysis owing to access limitations to the Advanced Light Source, but these 2 days represented different degrees of photochemical aging, the significance of which will be discussed below. Fifty-four samples from the intensive filter sampling periods, collected at a time resolution of 3-6 h, were analyzed using electron impact (EI) ionization. Sampling at the Bakersfield site was conducted at the southeast edge of the city, colocated with other instruments participating in the field campaign. Two 23 h samples, collected on Friday, 18 June, and Wednesday, 23 June, from midnight to 11 P.M. local time, were selected for VUV analysis. These two samples, along with 32 other daily (midnight to 11 P.M.) filter samples, were also analyzed using EI. All samples collected from the LA site were collected on quartz fiber filters (Tissuquartz[™] Filters, 2500 QAT-UP, Pall Life Sciences), which were $20 \text{ cm} \times 25 \text{ cm}$, allowing for high-volume PM_{2.5} sampling at $\sim 1 \text{ m}^3 \text{ min}^{-1}$. All samples collected from the Bakersfield site were collected on 86 cm^2 quartz fiber filters (Pall Life Sciences); however, medium-volume PM_{2.5} sampling at 226 L min^{-1} was conducted. All filters were prebaked at 550°C for 6 h to remove any organic contaminants. Filter samples were stored in a freezer at -18°C until chemical analysis. Field blanks were collected every 7–10 days by placing a prefired quartz fiber filter into the sampler for 15 min before removing and storing in the same manner as the field samples. Analysis of both field and laboratory blanks showed no significant aliphatic contaminants on prefired quartz filters.

[7] Additional samples collected during a study conducted at the Caldecott Tunnel in Oakland, California, were also analyzed for comparison [Dallmann et al., 2012]. Here we focus our analysis on two weekday filter samples collected in the gasoline-only bore and five samples collected in the mixed gasoline + diesel bore (four weekday and one weekend samples). Emissions during collection of the two weekday samples from the gasoline-only bore and the weekend sample from the mixed bore are expected to be dominated by gasoline vehicles (three "gasoline-dominated" samples), while emissions during the four weekday samples in the mixed bore are expected to be dominated by diesel vehicles (four "diesel-dominated" samples). In addition to thermal desorption of filter samples, SAE 10W-30 motor oil and three samples of crude oil extracted and refined in the San Joaquin Valley region were also analyzed for their chemical composition. The specific geographical origins of the crude oil samples are confidential for proprietary reasons.

2.2. Speciation of Semivolatile Hydrocarbons

[8] Selected samples were analyzed using comprehensive two-dimensional gas chromatography coupled to a vacuum ultraviolet high-resolution time-of-flight mass spectrometer (GC × GC/VUV-HRTOFMS). Filter punches (total area of 1.6 cm²) were thermally desorbed at 320°C under helium using a thermal desorption system and autosampler (TDS3 and TDSA2, Gerstel). Desorbed samples were focused at 20°C on a quartz wool liner in a cooled injection system (CIS4, Gerstel) before they were introduced into a gas chromatograph (GC, Agilent 7890) by rapidly heating the inlet to 320°C. For motor oil and crude oil, samples diluted in chloroform were directly injected via a septumless inlet into the GC.

[9] Comprehensive GC × GC was performed using a $60 \text{ m} \times 0.25 \text{ mm} \times 0.25 \text{ }\mu\text{m}$ nonpolar capillary column (Rxi-5Sil MS, Restek) for the first dimension separation (by volatility) and a medium-polarity second dimension column (1 m × 0.25 mm × 0.25 µm, Rtx-200MS, Restek). After sample injection, the GC oven was ramped from 40°C to 100°C at $60^{\circ}\text{C min}^{-1}$ and then 320°C at $3.5^{\circ}\text{C min}^{-1}$. A dual-stage thermal modulator (Zoex), consisting of a guard column (1 m × 0.25 mm, Rxi, Restek), was used as the interface between the two columns, where the effluent from the first column was cryogenically focused and periodically heated for rapid transfer into the second column. The modulation period was 2.4 s. The second dimension column was housed in a secondary oven, which was maintained at 15°C above the main oven temperature.

[10] Effluent from the second column was analyzed using a high-resolution ($m/\Delta m \sim 4000$) time-of-flight mass spectrometer (Tofwerk, Thun, Switzerland) coupled to the Advanced Light Source at Lawrence Berkeley National Laboratory. Single photon ionization by the vacuum ultraviolet (VUV) beam used here is similar to that reported previously [Isaacman et al., 2012b]. In brief, photons of 10.5 eV with an energy distribution width of 0.2 eV and an intensity of $\sim 10^{15}$ photons s⁻¹ were generated by the Chemical Dynamics Beamline (9.0.2). The photon beam was introduced orthogonally to the GC effluent, and ionized molecules were accelerated into the extraction region. Since the ionization energies of most organic compounds are between 8 and 11 eV, the minimal excess energy (compared to electron impact ionization at 70 eV) limits fragmentation of ionized molecules, allowing for significant detection of the molecular ions (M⁺⁺). The ion chamber was maintained at 150°C to minimize fragmentation. All LA and Bakersfield samples were also analyzed under EI at 70 eV, using a standard tungsten filament. The frequency of data collection was set at 200 Hz.

[11] The molecular ion signals for linear, branched, and cyclic alkanes under VUV ionization are used as the basis for quantification [Isaacman et al., 2012a]. Details about the calibration method based on molecular structures are described in Text S1 of the supporting information. The adjusted molecular ion signals were then quantified based on known quantities (5 ng each) of perdeuterated internal standards (n-hexadecane-d₃₄, n-eicosane-d₄₂, n-tetracosane-d₅₀, and n-octacosane-d₅₈, C/D/N Isotopes) spiked onto the filter punches prior to analysis. In quantifying the aliphatic hydrocarbons, only the UCM region of the chromatogram was considered (second dimension retention time between 0.4 and 0.9 s). Polycyclic aromatic hydrocarbons (PAHs) and highly oxygenated compounds are not expected to elute in this region. In addition, signals of oxygen-containing molecular ions were removed by high-mass-resolution data-processing procedures [Isaacman et al., 2012a], such that only aliphatic molecular ions (C_xH_v⁺) were considered for quantification. All data processing and visualization of GC × MS data were performed using custom code written in Igor 6.2.2 (Wavemetrics) adapted from high-resolution analysis of Aerodyne high-resolution time-of-flight mass spectrometer (AMS) data [DeCarlo et al., 2006], while visualization of two-dimensional chromatography data and detection of chromatographic peaks were carried out using GC Image software (LLC).

2.3. Other Measurements

[12] Colocated measurements of volatile organic compounds (VOCs) and OA are also used for comparison. Details of the GC/MS technique used to measure VOCs in LA are described by *Gilman et al.* [2010]. Average photochemical age of air masses at the LA site was calculated using the ratio of 1,2,4-trimethylbenzene to benzene [Parrish et al., 2007; Borbon et al., 2013]. GC/MS measurements of light hydrocarbons at the Bakersfield site were used for source identification purposes. Details of the GC/MS technique are described by Gentner et al. [2012]. Mass concentrations of organic aerosol were measured using an Aerodyne high-resolution time-of-flight mass spectrometer (hereafter referred to as AMS) [Hayes et al., 2013].

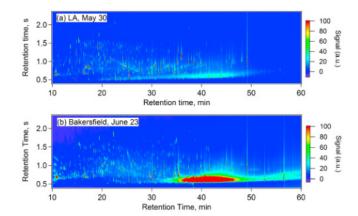


Figure 1. Total ion chromatogram of (a) LA sample (30 May) and (b) Bakersfield sample (23 June). The unresolved complex mixture (UCM) of aliphatic hydrocarbons, between first dimension retention times of 30 and 50 min, is evident in both samples.

3. Results and Discussion

3.1. Resolving the UCM Using GC × GC and GC × MS

[13] Figure 1 shows the total integrated ion chromatogram of the LA and Bakersfield samples under VUV ionization. Similar to previous work using GC × GC to analyze atmospheric samples, PAHs and oxygenated compounds, such as acids and ketones, are readily separated from aliphatic compounds [Worton et al., 2012]. Examples of these peaks are provided in Figure S2 in the supporting information. Despite the increased peak resolution, there is still a large unresolved "hump" in the nonpolar region of the chromatogram, representing the UCM. The dominance of the UCM is present in most ambient samples that were analyzed by EI. The inability of GC × GC to resolve these aliphatic hydrocarbons is likely due to the large number of possible constitutional isomers [Goldstein and Galbally, 2007], resulting in overlapping volatilities. In addition, even though the isomers are structurally different, all hydrocarbons are nonpolar and are essentially unretained in a polar column, resulting in poor second dimension resolution. Further chromatographic separation of the UCM can be performed using other column pairings [Nelson et al., 2006; Vogt et al., 2007] or coupled liquid chromatography-GC × GC [Mao et al., 2009]. However, in analyzing atmospheric samples, a polar column is preferred owing to the need to also separate aromatic and oxygenated compounds, which are always present. Orthogonal separation by volatility and polarity also allows for convenient inputs into two-dimensional frameworks used to model evolution of organic compounds in the atmosphere [Jimenez et al., 2009; Isaacman et al., 2011].

[14] In this work, the aliphatic UCM is characterized by a novel method using a combination of gas chromatography and soft VUV photoionization. Traditional GC/MS employs electron impact at 70 eV as the ionization technique, which imparts a large amount of excess energy and causes extensive fragmentation. While the fragmentation mass spectra can in many cases be useful for compound identification, aliphatic compounds have almost identical fragmentation patterns, with a large fraction of signal at m/z 41, 43, 55, 57, 69, 71, etc. (corresponding to $C_xH_{2x+1}^+$ and $C_xH_{2x-1}^+$ ions, where x is typically between 3 and 6). As a result, these compounds

cannot be distinguished from each other, as illustrated in Figure 2a. Soft ionization retains the identity of a molecule by maximizing the relative signal of the molecular ion, albeit at the loss of ionization efficiency by a factor of 20–70 [Northway et al., 2007]. Similar to aliphatic compounds in diesel fuel shown in our previous study, coupling soft ionization with volatility-based GC separation allows hydrocarbons in the UCM in the ambient atmosphere to fall into distinguishable patterns in a GC×MS plot (Figure 2b) [Wang et al., 2005; Isaacman et al., 2012b].

[15] In the GC×MS diagram (Figure 2), all compounds with the same carbon number fall into the same cluster. The carbon number corresponding to each cluster is identified using known retention times of n-alkanes and their molecular weights. Within each cluster, the molecular formulas of different isomers are identified from molecular weights of parent ions, shown on the y axis of Figure 2b. Every decrease of 2 in parent m/z (i.e., loss of 2 H atoms) corresponds to a double bond or a ring. As a result, the complex mixture can be speciated by the number of carbon atoms and the number of double bond equivalents, N_{DBE} , which is the total number of rings, double bonds, and twice the number of triple bonds. In the samples from both urban sites, the observed UCM lies in the range of 20-25 carbon atoms. In this range, alkanes have saturation vapor pressures between 9×10^{-8} and 9×10^{-6} Torr (saturation concentrations between approximately 2 and $200 \,\mu g \, m^{-3}$ at $298 \, K$) and are therefore semivolatile and expected to exist in both the particulate and gas phases under relevant ambient loadings [Williams et al., 2010a]. It must be noted that smaller, more volatile hydrocarbons are expected to be abundant but cannot be effectively and quantitatively collected on quartz filter samples used in this study. Most hydrocarbons observed have $N_{\rm DBE}$ < 6. Alkenes in this molecular weight range are highly reactive and are not expected to be significant in atmospheric samples. While alkene production from thermal decomposition of labile compounds (containing oxygen or nitrogen) cannot be ruled out, the distributions across different N_{DRE} classes were similar to those observed in unoxidized motor oil [Isaacman et al., 2012a]. Therefore, cyclic alkanes (not alkenes or alkynes) are expected to dominate aliphatic compounds observed with $N_{\rm DBE} > 0$ in Bakersfield and LA, similar to alkanes found in motor oil.

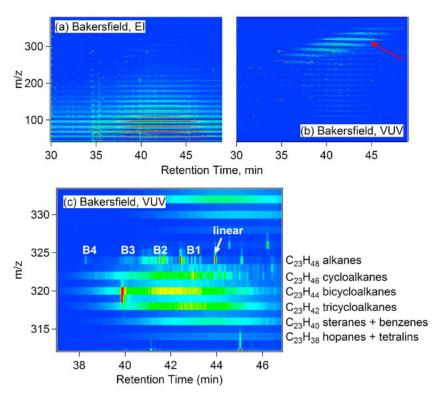


Figure 2. GC × MS plots of UCM in Bakersfield sample (23 June) under EI and VUV ionization. (a) Under electron impact, strong fragmentation causes most of the signal to fall on the same fragment $C_xH_y^+$ ions, as delineated by horizontal lines along a few select m/z. (b) Under VUV, molecular ions are retained and cluster by carbon number in a GC × MS plot. (c) One such cluster denoted by the red arrow is expanded. Within each carbon number, aliphatic hydrocarbons were resolved by number of rings (by molecular weight) and by degree of branching for acyclic alkanes, denoted by Bx, where x is the number of alkyl branches.

[16] Within each double bond equivalents class, gas chromatography provides separation of structural isomers (with the same molecular formula) by volatility. Since branched alkanes are more volatile than their linear counterparts of the same carbon number, they elute earlier from the first dimension column. As shown in Figure 2c, good chromatographic resolution is achieved for linear and branched alkanes with $N_{DBE}=0$. Unlike in diesel fuel [Isaacman et al., 2012b] but as observed in motor oil [Isaacman et al., 2012a], structural isomers are not well resolved for $N_{\text{DBE}} > 0$. This is most likely due to the higher carbon number range of the ambient UCM and hence, a greater number of possible geometric (cis/trans) isomers. As a result, mass concentrations of branched isomers are only reported for acyclic alkanes ($N_{\rm DBE}$ = 0). Following the naming convention of our previous work, we refer to branched isomers which contain x alkyl branches as Bx isomers. For example, B0 refers to the linear alkane, B1 refers to alkanes with one alkyl branch, etc.

[17] To the authors' knowledge, this work represents the most detailed characterization of the UCM in atmospheric samples to date. The calibrated masses of linear, branched, and cyclic alkanes for carbon numbers between 20 and 25 are shown in Figure 3. The relative ratios of total branched alkanes to linear alkanes (hereafter referred to as *B/N* ratio) vary greatly between the two urban sites, from ~1 in the 5 June LA sample to >10 in both Bakersfield samples, shown in Figure 4. Here we investigate how the sources

and atmospheric reactions of these SVOCs affect their relative abundances.

3.2. Processing of Semivolatile Hydrocarbons in Los Angeles, California

[18] In both LA samples analyzed under VUV photoionization, while the unresolved "hump" is visible in the chromatogram, aromatic and more polar compounds have higher total integrated ion signals than the UCM (Figure 1a). The most abundant isomer for each carbon number is the linear alkane, accounting for >30% of the total acyclic alkane mass. The fraction of linear alkane is similar to those found in diesel fuel [Isaacman et al., 2012b], albeit at smaller carbon numbers. The two samples analyzed using VUV represent two different periods in the field campaign. From 2 to 6 June, there was a steady buildup of pollutants due to periodic entrainment of the residual layer aloft, causing fresh emissions in the boundary layer to be mixed with aged pollutants that remained in the basin [Pollack et al., 2012]. On 30 May, no such buildup was observed. Owing to this difference in meteorological conditions, the integrated OH exposure, calculated from the ratios of 1,2,4-trimethylbenzene (1,2,4-TMB) to benzene [Borbon et al., 2013], was higher during collection of the 5 June sample $(4.8 \times 10^{10} \text{ molecules cm}^{-3} \text{ s or } 8.9 \text{ h at OH} = 1.5 \times 10^{6}$ molecules cm $^{-3}$) than that of the 30 May sample (2.6 × 10 10 molecules cm $^{-3}$ s or 4.8 h).

[19] The B/N ratios of C_{21} – C_{24} hydrocarbons in the 30 May sample are higher than those in the 5 June sample,

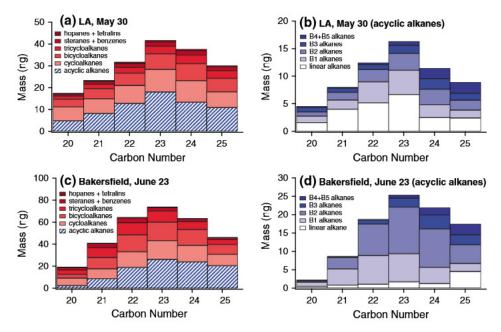


Figure 3. Distribution of alkane isomers in LA (30 May) and Bakersfield (23 June) samples. (a and c) Isomers are grouped by $N_{\rm DBE}$, ranging from 0 (acyclic) to 6 (hopanes and tetralins). (b and d) The signals include only those of acyclic (linear and branched) alkanes. While the distributions among different $N_{\rm DBE}$ in LA and Bakersfield are similar, the branched acyclic alkanes are more abundant relative to linear alkanes in Bakersfield than in LA.

suggesting that branched alkane isomers are consumed more rapidly relative to the linear isomers of the same carbon number. Since the aliphatic hydrocarbons in the UCM are predominantly saturated, their primary sink in the atmosphere is reaction with OH radicals. Although compounds in this volatility range exist in both gas and particle phases, C₂₁—C₂₄ hydrocarbons are sufficiently volatile that their gas-phase reactions with OH are dominant over particle-phase reactions at ambient levels of OA concentrations [Lambe et al., 2009]. We therefore expect that the change in B/N ratios is primarily a result of differences in gas-phase reaction rates with OH radicals, and the decay can be described by the following equation:

$$\left(\frac{B}{N}\right)_{\text{sample}} = \left(\frac{B}{N}\right)_{0} e^{(k_{N}f_{N} - k_{B}f_{B})} \int [OH] dt$$
 (1)

[20] $(B/N)_0$ is the B/N ratio at the source. $(B/N)_{\text{sample}}$ is the B/N ratio of each individual sample. k_B is the average OH reaction rate of all branched isomers (cm³ molecule $^{-1}$ s $^{-1}$), which is treated here as an unknown. k_N is the OH reaction rate of the linear alkane (cm³ molecule $^{-1}$ s $^{-1}$), calculated from structural-reactivity relationships [Kwok and Atkinson, 1995]. f_N is the average gas-phase fraction of the linear isomers. f_B is the average gas-phase fraction of branched isomers. f_B is the integrated OH exposure (molecules cm $^{-3}$ s) for that sample, calculated from the 1,2,4-TMB to benzene ratios [$Borbon\ et\ al.$, 2013]. Equation (1) is similar in form to previous investigation of hydrocarbon clocks for gas-phase oxidation, with one notable exception. Here we include the effect of gas-particle partitioning, which reduces the fractions of alkane species in the gas phase available for

OH reactions. The gas-phase fraction for each filter data point is calculated individually based on known vapor pressures [Pankow and Asher, 2008; Williams et al., 2010a], total OA mass concentrations measured by AMS [Hayes et al., 2013], and temperature. Both OA mass concentrations and temperatures were similar between the two filter sampling periods (at $7 \mu \text{g m}^{-3}$ and 290 K). For the data shown in Figure 5, we assume that f_N and f_B are constant to derive k_B for each carbon number.

[21] To investigate the relationship between B/N ratios and OH exposure, the B/N ratios at the source (corresponding to zero OH exposure) must be known. Since

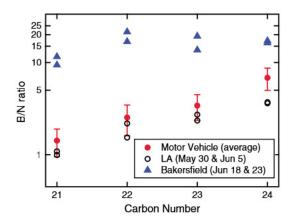


Figure 4. Ratios of branched to linear alkanes (*B/N*) of CalNex ambient samples and Caldecott tunnel samples (motor vehicle) analyzed using VUV photoionization. For the motor vehicle samples, the *B/N* ratios shown here are averaged over all seven samples, and the error bars represent the standard deviations in the seven samples.

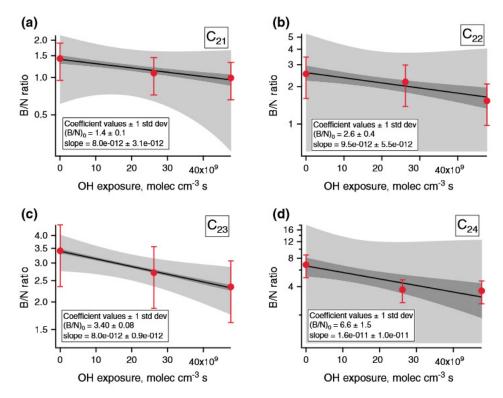


Figure 5. Ratios of branched to linear alkanes (B/N) of LA samples measured by VUV as a function of OH exposure. The B/N ratios at zero OH exposure are taken to be those from the tunnel samples of fresh motor vehicle exhaust. OH exposure for the LA samples was calculated using the ratio of 1,2,4-TMB to benzene [Borbon et al., 2013]. The black lines represent the best fit lines to equation (1), and the fitted coefficients are shown in Table 1. The dark grey and light grey shaded areas represent the 68% and 95% confidence intervals of the regressions, respectively.

the major contributor of hydrocarbons in the LA Basin is likely motor vehicles [Williams et al., 2010b], the B/N ratios of hydrocarbons in the LA ambient samples are compared to those in fresh vehicular exhaust, measured during a tunnel study in Oakland, California, in July 2010. Here we focus on the branched and linear acyclic alkanes between C21 and C24. The average B/N ratios of these alkanes in all seven tunnel samples are presented in Figure 4. There were no systematic differences in B/Nratios between diesel-dominated and gasoline-dominated samples in this carbon number range, so here we report the standard deviation in the seven samples to be the uncertainty in B/N ratios for each carbon number. The average B/N ratios observed in these tunnel samples, which are higher than those observed in the LA samples, serve as additional data points corresponding to zero OH

exposure at the source. Assuming that the decrease in B/N ratio is entirely due to photooxidation, the relative reaction rates of the branched isomers can be estimated from the decay of B/N ratios. A plot of B/N ratio versus OH exposure is shown in Figure 5. From regression using equation (2), a value for k_B for each carbon number can be derived from the slope of the best fit line $(f_Nk_N - f_Bk_B)$. In general, the average rate constants of branched alkane isomers are 21–35% higher than that of the linear alkanes (see Table 1).

[22] The uncertainties in the regressions on the *B/N* ratios measured using VUV are considerable, owing to the limited number of samples analyzed using this technique. To further constrain the dependence of *B/N* ratios on photochemical processing, analysis of samples covering a wide range of photochemical processing is needed. As mentioned before, using electron impact, detailed speciation, similar to that

Table 1. Average Gas-Phase OH Reaction Rate Constants of Linear (k_N) and Branched (k_B) Alkanes Calculated by Structure-Reactivity Relationships (SRR) [Kwok and Atkinson, 1995] and the Rate Ratios of B1 (k_{B1}/k_N) and B2 Isomers (k_{B2}/k_N) to Linear Isomers Calculated From SRR and Observed in Ambient LA Samples^a

Carbon Number	$k_N (10^{-11} \mathrm{cm}^3 \mathrm{molecule}^{-1} \mathrm{s}^{-1})$ From SRR	$k_B (10^{-11} \mathrm{cm}^3 \mathrm{molecule}^{-1} \mathrm{s}^{-1})$ From SRR	$\frac{k_{B1}/k_N}{({ m SRR})}$	k_{B2}/k_N (SRR)	k_B/k_N Observed (From VUV Data)	k_{B1}/k_N Observed (From EI Data)
21	2.66	2.88	1.08	1.01	1.21 ± 0.47	
22	2.80	3.03	1.08	1.01	1.31 ± 0.48	1.32
23	2.94	3.17	1.08	1.01	1.25 ± 0.13	1.12
24	3.08	3.31	1.07	1.01	1.35 ± 0.50	1.14

^aHere we assume B2 to contain two tertiary carbon atoms. The uncertainties in observed rate ratios are derived from the standard deviations in the regressions shown in Figure 5. See Text S3 in the supporting information for derivation of k_{B1} from EI data.

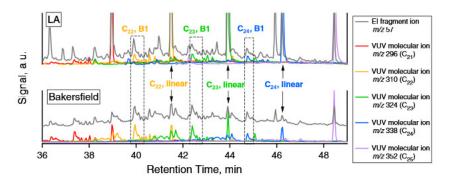


Figure 6. Extracted ion chromatograms of m/z 57 under EI (dark grey) and molecular ions of alkanes under VUV (all other colors). For each molecular ion, the latest eluting peak represents the linear alkane. Under EI, m/z 57 represents the $C_4H_9^+$ ion, the most abundant fragment ion of high molecular weight alkanes. The EI m/z 57 trace shows identical peak shapes to the molecular ion traces where the B1 isomers elute (denoted by boxes with dashed lines). Linear isomers (denoted by arrows) coelute with a B2 isomer of a higher carbon number, but for C_{23} – C_{24} in Bakersfield samples and C_{22} – C_{24} in the LA samples, the contributions of B2 isomers are expected to be less than 15%. The ratio of B1 to linear isomers (B1/N) for each carbon number can therefore be estimated from the m/z 57 signal under EI.

achieved in Figure 3, is difficult owing to extensive fragmentation and coelution of isomers. However, some B0 (linear) and B1 isomer peaks can be resolved from other isomers, and their retention times and molecular formulas are confirmed by VUV analysis, as shown in Figure 6. Each linear isomer coelutes with a B2 isomer with one more carbon atom, but based on observations from VUV, this B2 isomer is expected to contribute <15% of the total ion signal for C22-C24 alkanes in the LA samples. Other isomers, such as B2 isomers and isomers with $N_{\rm DBE} > 0$, remain unresolved using EI. The ratio of B1 to linear isomers (B1/N ratios) for each carbon number was calculated from the backgroundsubtracted signals of $C_4H_9^+$ (m/z 57), the most common alkane fragment ion. Using this method, more reliable estimates of k_B are derived from EI data, which consist of a larger number of samples. However, temperature and OA mass can no longer be assumed to be constant, and these factors critically determine gas-phase fractions and hence oxidation rates of SVOCs. For the EI data set, we first determine the average photochemical age for each sample, assuming an [OH] of 1.5×10^6 molecules cm⁻³. An average temperature and OA mass for that time period, to which each air plume represented by the sample has been exposed, are then determined. With this information, a B1/N ratio for each sample can be calculated for any given k_{B1} and $(B1/N)_0$ using equation (1), and optimal values for these two parameters are determined that minimizes the absolute error between calculated and observed B/N ratios derived from EI data. The regressions are shown in Figure S3 in the supporting information. Values for k_{B1} are consistent with those derived from VUV data, as shown in Table 1, confirming that B1 isomers of C22-C24 alkanes are consumed more rapidly than linear isomers. Figures 7a–7c show the diurnal profiles of the B1/N ratios, and the minima in B1/N ratios occur between late morning and early afternoon. The daytime minima are likely a result of higher OH concentrations and greater extent of oxidation. In a later section, we will rule out higher daytime temperatures (which alter gas/particle partitioning) to be the cause of the observed drop in B1/N ratios.

[23] Structure-reactivity predictions for linear and branched alkanes with 21-24 carbons are summarized in

Table 1. The predicted rate constants of branched alkanes are 1–8% higher than those of the linear isomers, lower than the ratios observed. Under these predictions, a combination of a tertiary and a primary carbon atom has higher OH reaction rate than two secondary carbon atoms if the tertiary carbon atom is not at the 2-position [Kwok and Atkinson, 1995]. While small (<C₅) branched hydrocarbons have lower OH rate constants than their linear isomers, it is more statistically likely for C21-C24 alkanes that alkyl branching will instead occur at the 3-position or higher positions, resulting in higher predicted OH rate constants. It is also noteworthy that preferred depletion of branched isomers was also observed during experimental studies of heterogeneous oxidation of motor oil particles [Isaacman et al., 2012a]. In those experiments, oxidation is expected to occur in the particle phase owing to the high particulate loadings, and the ratios of the rate constants $(k_B/k_N=1.28)$ reported are roughly consistent with those of gas-phase rate constants observed in this work.

[24] While the possibility of additional sources with low B/N ratios cannot be ignored, differences in this ratio caused by some potential primary sources of aliphatic hydrocarbons can be ruled out. There is no odd carbon preference in this range of carbon numbers, indicating that plant wax is not a major contributor to aliphatic compounds in the UCM. Also, wood burning markers, such as dehydroabietic acid and retene, were not observed at significant concentrations, indicating that the trend in B/ N ratios was not caused by emissions of linear alkanes from biomass or biofuel burning, and consistent with results from AMS analysis [Hayes et al., 2013]. No additional source of linear alkanes contributing to the lower B/ N ratios in more aged samples has yet been identified. We also considered whether the changes in B/N ratio could be caused by a difference in fleet between Oakland, California, and the LA Basin, but concluded that differences would likely be minimal based on comprehensive fuel analysis across the state [Gentner et al., 2012]. Additionally, this would only affect B/N ratios of the emission source $(B/N)_0$ and would be unlikely to cause B/N ratios to decrease with photochemical age as demonstrated.

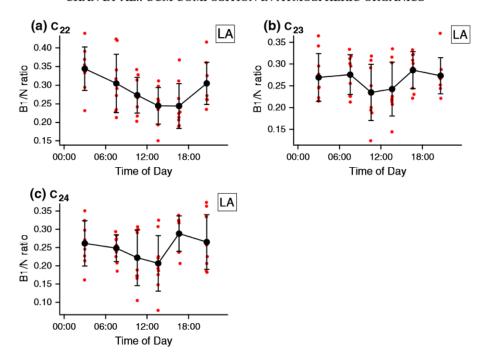


Figure 7. Diurnal profiles of B1/N ratios in LA. The minima of B1/N ratios occur between late morning and early afternoon local time, consistent with maximum photochemical activity.

3.3. Source of Semivolatile Hydrocarbons in Bakersfield

[25] Semivolatile and nonvolatile hydrocarbons in Bakersfield appear to originate from a broader combination of sources than in LA. First, the high concentrations of odd-number n-alkanes (C_{23} and C_{25} ; see Figure 3d) suggest that plant wax could be a major source. The carbon number preference index (CPI) of n-alkanes between C_{21} and C_{33} is given by the following equation [Marzi et al., 1993]:

$$CPI = \frac{\sum_{i=10}^{15} C_{2i+1} + \sum_{i=11}^{16} C_{2i+1}}{2\sum_{i=10}^{15} C_{2i}}$$
(2)

where C_x is the concentration of *n*-alkanes with *x* carbon atoms. The CPI of Bakersfield samples were consistently above 2, indicating a strong contribution of plant wax. In contrast, the average CPI of LA samples were below 1.4. This is consistent with factor analysis using AMS and Fourier transform infrared data from CalNex Bakersfield, which suggests that up to 10% of organic aerosol is likely vegetative detritus [Liu et al., 2012].

[26] Using GC/VUV, the alkane isomers in the UCM are classified into branched, cyclic, and linear alkanes based on molecular weights and GC retention times. Branched isomers dominate over linear isomers in Bakersfield, with B/N ratios greater than 10. These B/N ratios are inconsistent with those measured in the LA samples and are even higher than those observed in fresh motor vehicle exhaust from the tunnel samples. Since semivolatile branched alkanes react with OH radicals faster than linear isomers both theoretically (from structure-reactivity relationships) and in flow tube experiments [Isaacman et al., 2012a], photochemical processing is not expected to increase the B/N ratios from those of motor vehicle UCM to those observed in Bakersfield. While factor

analysis of particle-phase molecular markers suggests that a major source of linear alkanes is motor vehicles (Zhao et al., Sources of organic aerosol investigated using organic compounds as tracers measured during CalNex in Bakersfield, submitted to *Journal of Geophysical Research*, 2013), the high *B/N* ratio of the UCM suggests that there is also a large source of predominantly branched semivolatile hydrocarbons in addition to motor vehicles contributing significantly to the Bakersfield UCM.

[27] To investigate the temporal trends in this potentially large source of semivolatile branched hydrocarbons, the B1/N ratios are derived from EI data using the method previously described. B1/N ratios of only C23 and C24 can be reliably determined from EI data, owing to coelution of other linear alkanes with B2 isomers (see Figure 6). Extensive comparisons to speciated VOC concentrations in the region showed poor correlations of B1/N ratios to volatile tracers of known sources, such as gasoline and diesel vehicles, and natural gas production. One potential source of semivolatile hydrocarbons is the oil extraction and refining operations in the area, as Bakersfield is surrounded by numerous oil and gas fields, and two refineries are within 15 km of the sampling site. The isomer distributions of C₂₁-C₂₄ hydrocarbons of the ambient Bakersfield samples are therefore compared to those of three crude oil samples and one motor oil (SAE 10W-30) sample. As shown in Figure 8, the B/N ratios of the three crude oil samples vary by an order of magnitude. It is well known that the chemical composition of crude oil is highly variable, and the purpose of refining operations is to modify the chemical composition (relative concentrations of alkanes, alkenes, cycloalkanes, aromatics, etc.) for fuel, oil, chemical feedstock, or other uses. Lubricating oil, for example, has higher concentrations of cycloalkanes (also known as naphthenes) and branched alkanes than n-alkanes, as the viscosities of n-alkanes have

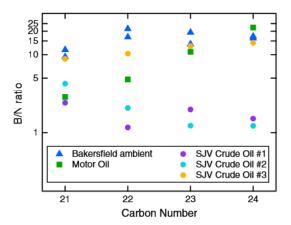


Figure 8. Ratios of branched to linear alkanes (*B/N*) of ambient Bakersfield samples, 10W-30 motor oil, and three samples of crude oil extracted and/or processed in the San Joaquin Valley (SJV). Among all the samples analyzed in this work, the *B/N* ratios of motor oil and one SJV crude oil sample (sample 3) approach those in Bakersfield, suggesting that the source of these hydrocarbons is related to oil operations in the area.

stronger temperature dependence and increase significantly when motor vehicle engines heat up [Gary et al., 2007]. While detailed characterization of the crude oil samples is beyond the scope of this work, the highly branched alkane content observed in Bakersfield is consistent with that in motor oil and one of the three crude oil samples analyzed. At this point, only qualitative comparisons are possible, but based on available data, the high concentrations of branched alkanes in the UCM are likely related to oil extraction and/or refining operations. In the future, detailed speciation of alkane isomers in direct emissions from oil operations could be useful in quantitatively determining their importance as a source of semivolatile hydrocarbons in this area.

3.4. Modeling Relative SOA Formation From Branched and Linear Alkanes

[28] Oxidation of semivolatile alkanes has been proposed to contribute significantly to urban SOA formation. While branched isomers are more likely to fragment upon oxidation [Lim and Ziemann, 2009], their emissions are higher and are oxidized more rapidly, owing to their higher volatility and greater reactivity toward OH radicals. Using the relative loss rates from CalNex LA, we formulate a simple box model to investigate the relative contributions to SOA formation from branched and linear alkanes. Based on equation (1), the ratio of SOA mass formed from branched alkanes to that from their linear isomer (hereafter referred to as $SOA_{B/N}$) for a particular carbon number can be described by the following:

$$SOA_{B/N} = \frac{SOA_B}{SOA_N} = \left(\frac{B}{N}\right)_0 \left(\frac{1 - e^{-k_B f_B \int [OH] dt}}{1 - e^{-k_N f_N \int [OH] dt}}\right) \left(\frac{Y_B}{Y_N}\right)$$
(3)

where Y_B and Y_N are the SOA mass yields for a branched alkane and a linear alkane, respectively. Here we calculate the theoretical SOA yields using a near-explicit gas-phase oxidation mechanism and theoretical partitioning based on work by *Jordan et al.* [2008]. We note that the SOA yields of

 C_{21} – C_{24} *n*-alkanes calculated from this approach are lower than those derived from extrapolating the volatility basis set parameters recommended by *Presto et al.* [2010] to this carbon number range. However, the approach by *Jordan et al.* [2008] is preferred here because the inclusion of a fragmentation mechanism is straightforward and the volatility basis set does not distinguish between different alkane isomers.

[29] The partitioning of alkanes between gas and aerosol phase $(f_B \text{ and } f_N)$ depends critically on organic loading and temperature. For our base case simulation, we assume an organic loading of $10 \,\mu g \, m^{-3}$ and an average temperature of 293 K, which is typical of ambient conditions during CalNex LA [Hayes et al., 2013]. To model the SOA yields Y_B and Y_N , one needs to know the fraction of peroxy radicals that forms alkoxy radicals and decompose into a smaller product, referred to as the "fragmentation branching ratio." Here we use 0.30 for this ratio in branched alkanes system. an intermediate value from laboratory studies [Lim and Ziemann, 2009], and assume it leads to a ketone with a carbon number half that of the parent alkane. For linear alkanes, no fragmentation is expected to occur in this carbon number range, and this fragmentation branching ratio is therefore zero. We assume that [OH] is constant at 1.5×10^6 molecules cm⁻³ [Hayes et al., 2013] and perform the simulation for 11 h, the maximum photochemical age calculated from observed 1,2,4-trimethylbenzene/benzene ratios. Figure 9 shows $SOA_{B/N}$ for this base case scenario. $SOA_{B/N}$ decreases with time because oxidation products of less reactive linear alkanes are formed later and partition into the particle phase

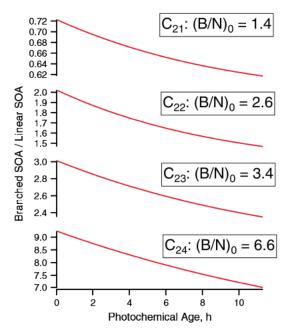


Figure 9. Base case simulation for simple box model. The mathematical equation is described by equation (3). For this simulation, the values for $(B/N)_0$ and k_B were derived from CalNex data summarized in Figure 5 and Table 1, respectively. SOA yields were calculated using the oxidation mechanism from *Jordan et al.* [2008], assuming 0.3 of reacted hydrocarbon fragments upon oxidation. The temperature of 293 K and organic loading is $10 \, \mu \mathrm{g \, m^{-3}}$, which are typical of conditions during CalNex LA. [OH] is assumed to be constant at 1.5×10^6 molecules cm⁻³.

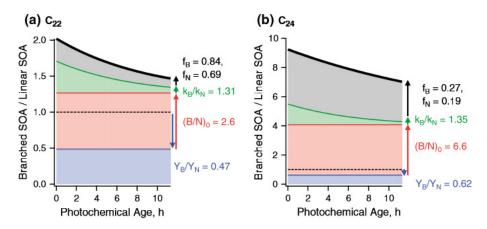


Figure 10. Effect of including various factors into determining relative SOA formation from (a) C_{22} and (b) C_{24} alkanes. The dotted line denotes equal SOA formation between all branched isomers and the linear alkane (SOA_{B/N}=1). The blue line and arrow denote lowering SOA yields of branched isomers as a result of fragmentation (0.3 in this base case). The red line and arrow denote including a higher emission ratio of branched alkanes, derived from motor vehicle samples (see Figure 5). The green line and arrow denote accounting for the differences in gas-phase oxidation rate constants, assuming all SVOCs are in the gas phase. Lastly, accounting for gas/particle partitioning results in the black line, which represents the base case scenario. The gas-phase fractions (f_N and f_B) are calculated for a temperature of 293 K and an organic loading of $10 \,\mu \mathrm{g} \,\mathrm{m}^{-3}$.

after those of branched alkanes do. However, as shown in the contour plots in Figure S4 and in equation (3), $SOA_{B/N}$ is linear with respect to $(B/N)_0$ and is therefore a stronger function of the emission ratios than photochemical age under typical urban conditions. As demonstrated in previous sections, $(B/N)_0$ can vary by an order of magnitude depending on the source. In the range of $(B/N)_0$ observed in this work, $SOA_{B/N}$ is expected to be at least 0.5 for all carbon numbers and can be greater than 10 in many cases. These results are qualitatively consistent with those from recent laboratory experiments and demonstrate that these previously unresolved alkanes are dominant SOA precursors, largely owing to greater abundances [Tkacik et al., 2012].

[30] Figure 10 shows the relative importance of different parameters that affect SOAB/N, including gas/particle partitioning (f_B and f_N), gas-phase kinetics (k_B and k_N), SOA yields, and emission ratios. Branched alkanes have lower SOA yields owing to greater extent of fragmentation. However, as a result of their higher volatilities, larger OH rate constants, and higher emissions, SOA formation from branched alkanes is expected to be 0.6 to 9 times that from linear alkanes in this carbon number range. In particular, the partitioning between gas and particle phases of the alkane precursors plays an important role in determining their availability for gas-phase OH reaction [Miracolo et al., 2010]. This effect is particularly apparent for C₂₄, where the difference in gas-phase fractions between linear and branched isomers is the greatest. It is therefore expected that temperature and organic loading have important effects on the relative oxidation rates through gas/particle partitioning and on subsequent SOA formation.

[31] To examine these effects individually, we carry out sensitivity analyses by performing the simulations at different temperatures and organic mass concentrations representing the typical range of conditions during CalNex. It is important to note that the effects of temperature and organic loading on oxidation rates are opposite to

those on SOA yields. Assuming gas/particle equilibrium, which was found to be a reasonable approximation for the LA site [Zhang et al., 2012], increasing organic mass leads to a smaller gas-phase fraction of both the alkane precursors and their oxidation products, leading to lower effective oxidation rates and higher SOA yields simultaneously. As shown in Figure 11a, the organic loading increases the relative contribution of branched SOA, indicating that these two effects are greater for branched alkanes than for linear alkanes. Figure 11b shows the effect of temperature on SOA_{B/N}. While increasing temperature increases fractions of both branched and linear alkanes in the gas phase (f_B and f_N , respectively), it increases f_N by a larger fraction than f_B , since branched alkanes are more volatile (i.e., $f_B > f_N$). Owing to this shift in partitioning at higher temperatures, the relative decay of branched to linear alkanes is expected to be smaller, so the relative contribution of branched isomers to SOA is reduced. This also rules out temperature fluctuations to be the cause of the observed daytime decrease in B/N ratios of the precursor alkanes shown in Figure 7, as higher daytime temperatures should lead to higher B/N ratios (of the precursor alkanes) and lower $SOA_{B/N}$. Here we examine the effect of temperature on partitioning of the alkanes only, since the partitioning parameters of their oxidation products (e.g., vapor pressures, heat of vaporization) and temperature dependence of SOA yields are poorly understood.

[32] Lastly, we used this box model to examine the effect of oxidation parameters on relative SOA formation (Figures 11c and 11d). As expected, an increase in oxidation rate constant of branched alkane relative to that of linear alkane translates into a linear increase in $SOA_{B/N}$. It should be noted that the photochemical age simulated here is typical of urban conditions (3 h at $[OH] = 1.5 \times 10^6$ molecules cm⁻³). On these time scales, the effects of relative kinetics are expected to be more dominant than on longer time scales. The relative amount of fragmentation versus functionalization also determines the

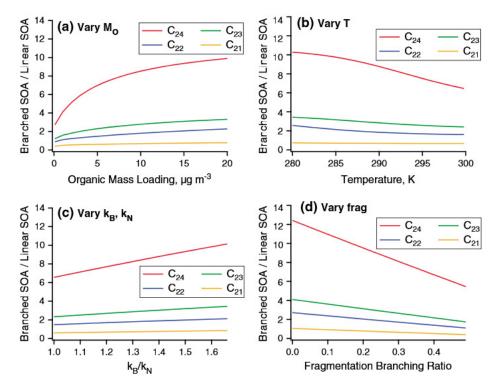


Figure 11. Effect of varying (a) organic loading, (b) temperature, (c) relative oxidation rate constants, and (d) fragmentation branching ratio on relative SOA formation. The range of organic mass loading and temperature simulated here is consistent with conditions during CalNex LA. The photochemical age for these simulations is 3 h, at an average [OH] of 1.5×10^6 molecules cm⁻³. Effects of temperature on SOA yields are not taken into account, as the heat of vaporization of oxidation products is poorly understood. The fragmentation branching ratio is the molar stoichiometric yield of the fragmentation product, which is assumed to be a ketone with half the carbon number as the parent alkane.

ratio of SOA yields, with fragmentation expected to be more significant for branched alkanes [Lim and Ziemann, 2009]. Here we show that the relative SOA formation varies linearly with changes in the fragmentation branching ratio. These results imply that the uncertainties in estimating k_B and k_N (~20% difference between structure-reactivity relationship (SRR) predictions and ambient measurements) have relatively small impacts on relative SOA contributions. On the other hand, more accurate estimates of branching between fragmentation and functionalization are needed to constrain SOA formation potentials from different alkane isomers.

4. Atmospheric Implications

[33] In this work, we present detailed characterization of SVOCs observed at two different urban sites to understand the sources and processing of semivolatile alkanes. In both LA and Bakersfield samples, GC×GC analysis provides excellent separation of polar and aromatic compounds. For the aliphatic UCM, GC coupled with soft photoionization mass spectrometry, such as VUV, provides detailed characterization of alkane isomers. The unresolved complex mixture of alkane isomers was resolved by carbon number, number of rings, and, for saturated alkanes, the degree of branching. This unprecedented level of detail provides interesting insights into the photochemical processing of semivolatile alkanes, which are important precursors to SOA in the atmosphere. The ratio of branched to linear

isomers for alkanes of the same carbon number decreased with photochemical processing in the LA Basin, consistent with a higher reaction rate of branched alkanes with OH radicals relative to linear alkanes. Detailed characterization of the UCM also provides information about sources, such as the odd carbon number preference and high relative concentrations of branched alkanes in the Bakersfield UCM consistent with the influences of plant wax and oil operations. Combining our speciated measurements (rate constants and emission ratios) and theoretical modeling (partitioning and SOA yields), branched alkanes are expected to contribute up to an order of magnitude more to SOA formation than linear alkanes, despite lower SOA yields. The relative contributions depend strongly on their emission ratios, branching ratios between fragmentation and functionalization, and gas/particle partitioning for this range of volatility.

[34] These novel measurements provide a basis to understand atmospheric processing of different hydrocarbons in the UCM, especially those from fossil fuel-related sources. The ability to speciate aliphatic compounds provides important constraints on both their effective oxidation rates and potentials to form SOA. As shown in this work, branched alkanes have higher OH reaction rates and their emissions from fossil fuel-related sources can be up to an order of magnitude higher than those of linear alkanes. Since branched alkanes react more rapidly with OH radicals than linear alkanes, the ratio of their concentrations serves as an indicator for degree of processing of SVOCs. This concept is similar to using

VOC ratios to estimate the gas-phase photochemical age but is more relevant to understanding SOA formation from oxidation of SVOCs. Our simple box model demonstrates that the oxidation rates of SVOC depend strongly on the fraction in the gas phase available for OH reaction [Miracolo et al., 2010]. As a result, parameters which affect partitioning, such as organic loading and temperature, also influence the oxidation kinetics. Quantitative measurements of gas/particle partitioning are therefore crucial to understanding the oxidation of SVOCs, which is expected to be their dominant loss process in the atmosphere. Also, when using SVOCs as source tracers or to estimate photochemical age, their gas/particle partitioning must be taken into account to accurately determine the extent of their atmospheric processing.

[35] Owing to differences in gas-phase oxidation chemistry, SOA yields depend strongly on molecular structure. Branched alkanes have a higher tendency to fragment upon oxidation, leading to more volatile products and lower SOA formation [Lim and Ziemann, 2009]. However, as demonstrated by our box model, oxidation of C₂₁-C₂₄ branched alkanes is expected to play a more important role in SOA formation than linear alkanes of the same carbon number, as a result of higher emissions, faster gas-phase oxidation rate constants, and higher volatility. The GC/VUV technique should be expanded to alkanes with fewer than 20 carbons, which comprise the intermediate volatility organic compounds, or IVOCs. A recent modeling study predicts that branched and linear IVOCs contribute a comparable compound of SOA [Pye and Pouliot, 2012], and future work using speciated GC/VUV measurements can be used to constrain their emissions, oxidation rates, and SOA yields.

[36] In this work, we focus our analysis on saturated alkanes. Experimental work has shown that the SOA yields of cyclic alkanes can be higher or lower than linear alkanes, depending on the degree of ring strain [Lim and Ziemann, 2009]. Here we show that their observed concentrations are greater than those of saturated alkanes. Their contributions to urban SOA formation are therefore potentially important, and accurate modeling will require a combination of speciated measurements using GC/VUV and laboratory studies into their oxidation chemistry and SOA formation mechanisms [Lim and Ziemann, 2009; Tkacik et al., 2012].

[37] Since SOA yields depend strongly on molecular structure, classification of organic species by volatility alone fails to fully capture known variations in SOA formation. Speciation by carbon number, number of rings, and alkyl branching shown in this work provides detailed knowledge of both volatility and molecular structure and hence insights into SOA formation from the UCM, the major component in many SVOC sources. More detailed chemical characterization of SVOCs may also give crucial information about their sources. With more time-resolved samples, measurements of compounds speciated by molecular structure can be more broadly applied in factor analysis and can serve as valuable inputs for source attribution. Ultimately, despite the complexity of organic mixtures, detailed speciation is an important direction providing essential understanding of emission sources and their environmental fates.

[38] Acknowledgments. This research was supported by the National Oceanic and Atmospheric Administration under award NA10OAR4310104.

The Advanced Light Source as well as K.R.W. and T.N. were supported by the Director, Office of Energy Research, Office of Basic Energy Sciences, of the U.S. Department of Energy under contract DE-AC02-05CH11231. Measurements at the Advanced Light Source were also supported by the Laboratory Directed Research and Development Program of Lawrence Berkeley National Laboratory under U.S. Department of Energy contract DE-AC02-05CH11231. Caldecott tunnel measurements were supported by EPA grant RD834553. The U.S. Environmental Protection Agency through its Office of Research and Development funded and collaborated in the research described here under contract EP-D-10-070 to Alion Science and Technology. The manuscript has been subjected to external peer review and has been cleared for publication. Mention of trade names or commercial products does not constitute endorsement or recommendation for use. P.L.H. and J. L.J. thank CARB 08-319/11-305, DOE (BER/ASR) DE-SC0006035, and a CIRES Visiting Fellowship to P.L.H. The authors would like to thank Sally Newman for use of temperature data.

References

Borbon, A., et al. (2013), Emission ratios of anthropogenic VOC in northern mid-latitude megacities: Observations vs. emission inventories in Los Angeles and Paris, J. Geophys. Res. Atmos., 118, 2041–2057, doi:10.1002/jgrd.50059.

Dallmann, T. R., S. J. Demartini, T. Kirchstetter, S. C. Herndon, T. Onasch, E. C. Wood, and R. A. Harley (2012), On-road measurement of gas and particle phase pollutant emission factors for individual heavy-duty diesel trucks, *Environ. Sci. Technol.*, 46(15), 8511–8518.

DeCarlo, P. F., et al. (2006), Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, Anal. Chem., 78(24), 8281–8289.

Frysinger, G. S., R. B. Gaines, L. Xu, and C. M. Reddy (2003), Resolving the unresolved complex mixture in petroleum-contaminated sediments, *Environ. Sci. Technol.*, 37(8), 1653–1662.

Gary, J. H., G. E. Handwerk, and M. J. Kaiser (2007), Petroleum Refining: Technology and Economics, 5th ed., Taylor and Francis, Boca Raton, Fla. Gentner, D. R., et al. (2012), Elucidating secondary organic aerosol from diesel and gasoline vehicles through detailed characterization of organic carbon emissions, Proc. Natl. Acad. Sci. U. S. A., 109(45), 18,318–18,323, doi:10.1073/pnas.1212272109.

Gilman, J. B., et al. (2010), Ozone variability and halogen oxidation within the Arctic and sub-Arctic springtime boundary layer, *Atmos. Chem. Phys.*, 10(21), 10,223–10,236, doi:10.5194/acp-10-10223-2010.

Goldstein, A. H., and I. E. Galbally (2007), Known and unexplored organic constituents in the Earth's atmosphere, *Environ. Sci. Technol.*, 41(5), 1514–1521.

Grieshop, A. P., M. A. Miracolo, N. M. Donahue, and A. L. Robinson (2009), Constraining the volatility distribution and gas-particle partitioning of combustion aerosols using isothermal dilution and thermodenuder measurements, *Environ. Sci. Technol.*, 43(13), 4750–4756.

Hallquist, M., et al. (2009), The formation, properties and impact of secondary organic aerosol: Current and emerging issues, Atmos. Chem. Phys., 9, 5155–5236, doi:10.5194/acp-9-5155-2009.

Hayes, P. L., et al. (2013), Aerosol composition and sources in Los Angeles during the 2010 CalNex campaign, J. Geophys. Res., doi:10.1002/ jgrd.50530, in press.

Hodzic, A., J. L. Jimenez, S. Madronich, M. R. Canagaratna, P. F. DeCarlo, L. Kleinman, and J. Fast (2010), Modeling organic aerosols in a megacity: Potential contribution of semi-volatile and intermediate volatility primary organic compounds to secondary organic aerosol formation, *Atmos. Chem. Phys.*, 10(12), 5491–5514, doi:10.5194/acp-10-5491-2010.

Isaacman, G., D. R. Worton, N. M. Kreisberg, C. J. Hennigan, A. P. Teng, S. V. Hering, A. L. Robinson, N. M. Donahue, and A. H. Goldstein (2011), Understanding evolution of product composition and volatility distribution through in-situ GC × GC analysis: A case study of longifolene ozonolysis, Atmos. Chem. Phys., 11(11), 5335–5346, doi:10.5194/acp-11-5335-2011.

Isaacman, G., A. W. H. Chan, T. Nah, D. R. Worton, C. R. Ruehl, K. R. Wilson, and A. H. Goldstein (2012a), Heterogeneous OH oxidation of motor oil particles causes selective depletion of branched and less cyclic hydrocarbons, *Environ. Sci. Technol.*, 46, 10,632–10,640.

Isaacman, G., et al. (2012b), Improved resolution of hydrocarbon structures and constitutional isomers in complex mixtures using gas chromatography-vacuum ultraviolet-mass spectrometry, *Anal. Chem.*, 84(5), 2335–2342, doi:10.1021/ac2030464.

Jimenez, J. L., et al. (2009), Evolution of organic aerosols in the atmosphere, Science, 326(5959), 1525–1529, doi:10.1126/science.1180353.

Jordan, C. E., P. J. Ziemann, R. J. Griffin, Y. B. Lim, R. Atkinson, and J. Arey (2008), Modeling SOA formation from OH reactions with C8– C17 n-alkanes, Atmos. Environ., 42(34), 8015–8026, doi:10.1016/j. atmosenv.2008.06.017.

- Kwok, E. S. C., and R. Atkinson (1995), Estimation of hydroxyl radical reaction rate constants for gas-phase organic compounds using a structure-reactivity relationship: An update, Atmos. Environ., 29(14), 1685–1695.
- Lambe, A. T., M. A. Miracolo, C. J. Hennigan, A. L. Robinson, and N. M. Donahue (2009), Effective rate constants and uptake organic molecular markers in motor oil and diesel primary radicals, *Environ. Sci. Technol.*, 43(23), 8794–8800.
- Lim, Y. B., and P. J. Ziemann (2009), Effects of molecular structure on aerosol yields from OH radical-initiated reactions of linear, branched, and cyclic alkanes in the presence of NO₃, Environ. Sci. Technol., 43(7), 2328–2334.
- Liu, S., et al. (2012), Secondary organic aerosol formation from fossil fuel sources contribute majority of summertime organic mass at Bakersfield, J. Geophys. Res., 117, D00V26, doi:10.1029/2012JD018170.
- Mao, D., H. Van De Weghe, R. Lookman, G. Vanermen, N. De Brucker, and L. Diels (2009), Resolving the unresolved complex mixture in motor oils using high-performance liquid chromatography followed by comprehensive two-dimensional gas chromatography, *Fuel*, 88(2), 312–318, doi:10.1016/j.fuel.2008.08.021.
- Marzi, R., B. E. Torkelson, and R. K. Olson (1993), A revised carbon preference index, Org. Geochem., 20(8), 1303–1306.
- Miracolo, M., A. Presto, A. T. Lambe, C. J. Hennigan, N. M. Donahue, J. H. Kroll, D. R. Worsnop, and A. L. Robinson (2010), Photo-oxidation of low-volatility organics found in motor vehicle emissions: Production and chemical evolution of organic aerosol mass, *Environ. Sci. Technol.*, 44, 1638–1643.
- Nelson, R. K., B. M. Kile, D. L. Plata, S. P. Sylva, L. Xu, C. M. Reddy, R. B. Gaines, G. S. Frysinger, and S. E. Reichenbach (2006), Tracking the weathering of an oil spill with comprehensive two-dimensional gas chromatography, *Environ. Forensics*, 7(1), 33–44, doi:10.1080/15275920500506758.
- Northway, M. J., et al. (2007), Demonstration of a VUV lamp photoionization source for improved organic speciation in an aerosol mass spectrometer, *Aerosol Sci. Technol.*, 41(9), 828–839, doi:10.1080/02786820701496587.
- Pankow, J., and W. Asher (2008), SIMPOL. 1: A simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, *Atmos. Chem. Phys.*, 8, 2773–2796, doi:10.5194/acp-8-2773-2008.
- Parrish, D. D., A. Stohl, C. Forster, E. L. Atlas, D. R. Blake, P. D. Goldan, W. C. Kuster, and J. A. de Gouw (2007), Effects of mixing on evolution of hydrocarbon ratios in the troposphere, *J. Geophys. Res.*, 112, D10S34, doi:10.1029/2006JD007583.
- Pollack, I. B., et al. (2012), Airborne and ground-based observations of a weekend effect in ozone, precursors, and oxidation products in the California South Coast Air Basin, J. Geophys. Res., 117, D00V05, doi:10.1029/2011JD016772.
- Presto, A. A., M. A. Miracolo, N. M. Donahue, and A. L. Robinson (2010), Secondary organic aerosol formation from high-NO_x photo-oxidation of low volatility precursors: n-Alkanes, Environ. Sci. Technol., 44, 2029–2034.
- Presto, A. A., C. J. Hennigan, N. T. Nguyen, and A. L. Robinson (2012), Determination of volatility distributions of primary organic aerosol emissions from internal combustion engines using thermal desorption gas chromatography mass spectrometry, *Aerosol Sci. Technol.*, 46(10), 1129–1139, doi:10.1080/02786826.2012.700430.
- Pye, H. O. T., and G. A. Pouliot (2012), Modeling the role of alkanes, polycyclic aromatic hydrocarbons, and their oligomers in secondary organic aerosol formation, *Environ. Sci. Technol.*, 46(11), 6041–6047, doi:10.1021/es300409w.
- Pye, H. O. T., and J. H. Seinfeld (2010), A global perspective on aerosol from low-volatility organic compounds, *Atmos. Chem. Phys.*, 10(9), 4377–4401, doi:10.5194/acp-10-4377-2010.

- Robinson, A. L., N. M. Donahue, M. K. Shrivastava, E. A. Weitkamp, A. M. Sage, A. P. Grieshop, T. E. Lane, J. R. Pierce, and S. N. Pandis (2007), Rethinking organic aerosols: Semivolatile emissions and photochemical aging, *Science*, 315(5816), 1259–1262, doi:10.1126/ science.1133061.
- Ryerson, T. B., et al. (2013), The 2010 California Research at the Nexus of Air Quality and Climate Change (CalNex) field study, J. Geophys. Res. Atmos., doi:10.1002/jgrd.50331, in press.
- Schauer, J., M. Kleeman, and G. Cass (1999), Measurement of emissions from air pollution sources. 2. C₁ through C₃₀ organic compounds from medium duty diesel trucks, *Environ. Sci. Technol.*, 33(10), 1578–1587.
- Schauer, J., M. Kleeman, and G. Cass (2002), Measurement of emissions from air pollution sources. 5. C₁-C₃₂ organic compounds from gasolinepowered motor vehicles, *Environ. Sci. Technol.*, 36, 1169–1180.
- Tkacik, D. S., A. A. Presto, N. M. Donahue, and A. L. Robinson (2012), Secondary organic aerosol formation from intermediate-volatility organic compounds: Cyclic, linear, and branched alkanes, *Environ. Sci. Technol.*, 46(16), 8773–8781, doi:10.1021/es301112c.
- Van Deursen, M., J. Beens, J. Reijenga, P. Lipman, and C. Cramers (2000), Group-type identification of oil samples using comprehensive two-dimensional gas chromatography coupled to a time-of-flight mass spectrometer (GC × GC-TOF), J. High Resolut. Chrom., 23, 507–510.
- Ventura, G. T., F. Kenig, C. M. Reddy, G. S. Frysinger, R. K. Nelson, B. Van Mooy, and R. B. Gaines (2008), Analysis of unresolved complex mixtures of hydrocarbons extracted from Late Archean sediments by comprehensive two-dimensional gas chromatography (GC × GC), Org. Geochem., 39(7), 846–867, doi:10.1016/j.orggeochem.2008.03.006.
- Vogt, L., T. Gröger, and R. Zimmermann (2007), Automated compound classification for ambient aerosol sample separations using comprehensive two-dimensional gas chromatography-time-of-flight mass spectrometry, *J. Chromatogr. A*, 1150(1-2), 2-12, doi:10.1016/j.chroma.2007.03.006.
- Wang, F. C.-Y., K. Qian, and L. A. Green (2005), GC x MS of diesel: A twodimensional separation approach, Anal. Chem., 77(9), 2777–2785, doi:10.1021/ac0401624.
- Williams, B. J., A. H. Goldstein, N. M. Kreisberg, and S. V Hering (2010a), In situ measurements of gas/particle-phase transitions for atmospheric semivolatile organic compounds, *Proc. Natl. Acad. Sci. U. S. A.*, 107(15), 6676–6681, doi:10.1073/pnas.0911858107.
- Williams, B. J., A. H. Goldstein, N. M. Kreisberg, S. V. Hering, D. R. Worsnop, I. M. Ulbrich, K. S. Docherty, and J. L. Jimenez (2010b), Major components of atmospheric organic aerosol in southern California as determined by hourly measurements of source marker compounds, Atmos. Chem. Phys., 10(23), 11,577–11,603, doi:10.5194/acp-10-11577-2010.
- Worton, D. R., N. M. Kreisberg, G. Isaacman, A. P. Teng, C. McNeish, T. Górecki, S. V. Hering, and A. H. Goldstein (2012), Thermal desorption comprehensive two-dimensional gas chromatography: An improved instrument for in-situ speciated measurements of organic aerosols, *Aerosol Sci. Technol.*, 46(4), 380–393, doi:10.1080/02786826.2011.634452.
- Zhang, Q., et al. (2007), Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, Geophys. Res. Lett., 34, L13801, doi:10.1029/2007GL029979.
- Zhang, X., J. Liu, E. T. Parker, P. L. Hayes, J. L. Jimenez, J. A. De Gouw, J. H. Flynn, N. Grossberg, B. L. Lefer, and R. J. Weber (2012), On the gas-particle partitioning of soluble organic aerosol in two urban atmospheres with contrasting emissions: 1. Bulk water-soluble organic carbon, J. Geophys. Res., 117, D00V16, doi:10.1029/2012JD017908.